Hydrothermal alterations mapping using Quickbird and Landsat-8 data, a case study from Babbiduyeh, Kerman province, Iran

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Abstract
In this work, we focus on investigating the Quickbird and Landsat-8 datasets for mapping hydrothermal and gossans alterations in reconnaissance porphyry copper mineralization in the Babbiduyeh area. This area is situated in the Central Iranian Volcano-sedimentary Complex, where large copper deposits like Sarcheshmeh as well as numerous occurrences of copper exist. The alteration zones are discriminated by implementation of band ratio and principal component analysis on the Quickbird and Landsat-8 datasets. The image processing results are evaluated by field surveys, X-ray diffraction (XRD), microscopic thin section, and spectroscopic studies of field samples as well as the 1:100000 Sarduiyeh and 1:5000 Babbiduyeh geological maps. In addition, the spectral characterizations of the samples are analyzed by visual inspection, and the PIMAView, SAMS, and ViewSpecpro software programs. The combined spectroscopic measurements, XRD analyses, and petrographic studies revealed mineral assemblages typical of the phyllic, phyllic-supergen, propylitic, argillic, and gossan alterations. The results obtained from image processing and analysis of field samples illustrated examples of effects of iron oxides and hydroxides on the surface of phyllic and argillic alterations. Hence, it can be concluded that the areas discriminated in Quickbird as gossans correspond to the phyllic and argillic alteration areas.

Keywords: Band Ratio, Hydrothermal Alteration, Landsat-8, Principal Component Analysis, Quickbird.

1. Introduction
Central Iranian Cenozoic Magmatic Belt (CICMB) has a significant potential for porphyry copper mineralization. Many researchers have applied the remote sensing and geographic information system techniques on CICMB for mapping hydrothermal alteration, especially in porphyry copper deposits [1-6]. Porphyry copper deposits are associated with the hydrothermal alteration zones such as phyllic, argillic, potassic, and propylitic. These zones may be affected by leaching and supergene alteration over porphyry copper deposits, and produce an oxide zone with hydroxide minerals including dissolution boxes, named gossan. These alteration zones contain specific minerals such as muscovite, kaolinite, biotite, alunite, chlorite, hematite, and jarosite, which show diagnostic absorption features in different parts of the electromagnetic spectrum, especially in the VNIR and SWIR region due to the chemical and mineralogical variations [7-10]. This characteristic offers a potential for mineral identification (especially, gold and copper) through remote sensing studies. The ability of remote sensing in discrimination of porphyry copper alterations using various sensors such as ASTER, TM, ETM+, and hyperion has been reported in many studies [1-5, 9-17]. However, relatively little publications are available for mapping alteration minerals using sensors such as OLI in Landsat-8 and Quickbird (a high spatial resolution sensor). Therefore, much more works should be carried out using these sensors. Quickbird that contains high spatial resolutions of
2.44 m and 0.61 m with four spectral bands in the VNIR region (covering the 0.45–0.90 μm range) may be suitable for discrimination of gossans (Table 1). The Quickbird dataset has been used in geological mapping and hazard assessing in the recent years [18-22]. Girouard et al. (2004) have compare the results between Quickbird and Landsat TM as high and medium spatial resolution sensors, respectively, to validate the Spectral Angle Mapper (SAM) algorithm for geological mapping in Central Jebilet Morocco [19]. Aydal et al. (2008) have used GIS and different processing methods including band combination, principal component analysis, crosta technique, band rationing, decorrelation stretching, and various filtering techniques on the ETM+ and Quickbird data to determine unknown ore deposits/enrichments in Hinzar Mountain in Turkey [20]. Liu et al. (2013) have used ETM+ and Quickbird for mapping lithological units, small intrusions, and alteration zones to target mineral resources in the Xiemisitai area, west Junggar, Xinjiang, China [21]. Mishra et al. (2014) have used the ASTER and Quickbird data to identify the chloritic and phyllic alteration zones and gossans related to gold mineralization in North Sudan [22].

Landsat-8 satellite is a new product from NASA under Landsat open source series, which was launched in February 2013. It carries two sensors including the Operational Land Imager (OLI) and Thermal Infrared Sensor (TIRS) along with three new bands: a deep blue band for coastal/aerosol studies, a shortwave IR band for cirrus detection, and a quality assessment band. The Landsat-8 data can detect the alterations and iron oxyhydroxide minerals due to the absorption and reflectance characteristics of these minerals through the OLI (Operational Land Imager) sensor with two spectral bands in the short-wave infrared (SWIR) and 5 bands in the visible near infrared (VNIR) (Table 2). Since the advent of Landsat-8, many researchers have used the landat-8 data in their studies [24-29]. Fotovat Roudsari and Solgi (2014) have used the Landsat-8 and ASTER data for identification of altered tuff and supergene zone to prospect polymetallic mineralization in Ferdows area, Iran [30]. Ahmed and Beiranvand Pour (2014) have used the conventional image processing methods such as color composite, principle component analysis, band ratio, and minimum noise fraction on the Landsat-8 data to find and map the lithological units and alteration zones in a gold mining area in NE Sudan [28]. Khalid et al. (2014) have carried out general prospecting in the gold deposit near Khartoum, Sudan, and searched for possible mineral deposits through an integrated analysis of the satellite gravity and Landsat 8 OLI data [31]. Beiranvand Pour and Hashim (2014) have studied the applicability of Landsat-8 for mapping hydrothermal alteration areas and lithological units associated with porphyry copper exploration in Sarcheshmeh Copper Mine, SE Iran (29). Mwaniki et al. (2015) have compared Landsat-8 and 7 for geological mapping and visualizing lineaments in the central region of Kenya [32]. The mentioned studies revealed the importance of Landsat-8 and Quickbird in geological studies and mineral mapping.

The main purpose of this work was to investigate the Quickbird and Landsat-8 datasets to map hydrothermal alterations and gossan in reconnaissance porphyry copper mineralization in the Babbiduyeh area. The other purpose was to validate and find the relationship between these two satellite images with field studies and sample analysis such as X-ray diffraction (XRD) analyzes, microscopic thin sections, and spectroscopy studies.

<table>
<thead>
<tr>
<th>Table 1. Characteristics of Quickbird sensor [23].</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Launch Date</strong></td>
</tr>
<tr>
<td><strong>Orbit Altitude</strong></td>
</tr>
<tr>
<td><strong>Swath Width</strong></td>
</tr>
<tr>
<td><strong>Digitization</strong></td>
</tr>
<tr>
<td><strong>Resolution</strong></td>
</tr>
<tr>
<td></td>
</tr>
<tr>
<td><strong>Image Bands</strong></td>
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<tr>
<td></td>
</tr>
</tbody>
</table>

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Table 2. Characterizations of Landsat-8 sensor [33].

<table>
<thead>
<tr>
<th>Band Name</th>
<th>Wavelength</th>
<th>Spatial Resolution</th>
</tr>
</thead>
<tbody>
<tr>
<td>Coastal/Aerosol</td>
<td>0.433-0.453 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Blue</td>
<td>0.450-0.515 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Green</td>
<td>0.525-0.600 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Red</td>
<td>0.630-0.680 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Near IR</td>
<td>0.845-0.885 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Shortwave IR</td>
<td>1.560-1.660 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Shortwave IR</td>
<td>2.100-2.300 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Panchromatic</td>
<td>0.500-0.680 µm</td>
<td>15 m</td>
</tr>
<tr>
<td>Cirrus</td>
<td>1.360-1.390 µm</td>
<td>30 m</td>
</tr>
<tr>
<td>Thermal IR</td>
<td>10.60-11.19 µm</td>
<td>100 m</td>
</tr>
<tr>
<td>Thermal IR</td>
<td>11.50-12.51 µm</td>
<td>100 m</td>
</tr>
</tbody>
</table>

2. Location and geological settings of studied area

The Babbiduiyeh area is located in the southern part of the Central Iranian Volcano-Sedimentary Complex (Uromiyeh-Dokhtar Belt) in the Kerman Cenozoic Magmatic Arc (KCMA) or Dehaj-Sarduiyeh copper belt, SE of Kerman province (Figure 1). The studied area is located in a mountainous semi-arid environment with low vegetation, in which the altitude varies between 1600 and 4391 m above the sea level, and therefore, it is ideal for remote sensing studies. CICMB is considered to have a suitable economic potential regarding copper mineralization, and it is analog to the important porphyry copper belts of the world such as Andean volcanic belt [34, 35]. Numerous occurrences of copper mineralization including the great deposit of Sarcheshmeh have been discovered at this belt.

Figure 1. A) Geographical location of studied area in Iran B) in Kerman Cenozoic Magmatic Arc (Dehaj-Sarduiyeh belt) [29] C) 1:5000 Geological map of Babbiduyeh [36].
Copper mineralization in the Babbiduyeh area is related to the Oligomiocene granodioritic intrusions, which penetrated in the volcano-sedimentary rocks of Eocene. According to the geological map of the Babbiduyeh area, the dominant lithology is igneous rocks. Dacites, riodacites, rhyolite, granite, and granodiorites in the area are related to Cenozoic Era, Neogen period. Other existent lithological units in the studied area involve andesite, andesite-basalt, and basalt with pyroclastic. In addition, conglomerate, trace sandstone, gravel, and recent age alluvium spread in the northern part of the region (Figure 1). Mineralization, which mostly occurs as porphyry copper, focuses at the central and southern part of Babbiduyed in sub-volcanic stock. Porphyry copper in Babbiduyeh is associated with the hydrothermal alteration zones such as phyllic, argillic, potassic, and propylitic. These zones were affected by leaching and supergene alteration, and produced an oxide zone with oxide and hydroxide minerals including dissolution boxes, named gossans (Figure 2).

3. Methods
The Quickbird and Landsat-8 images were used in this research work for mapping the alteration minerals. DigitalGlobe’s QuickBird image data was typically distributed in a relative radiance. Pre-processing was implemented on the data in order to remove noise and acquire surface reflectance. The QuickBird radiance calibration utility from ENVI 5.1 software was used to convert the relative radiance into the absolute radiance in units of [\(\mu W/(cm^2*nm*sr)\)]. The calibration was performed using the calibration factors in the QuickBird metadata file (the absCalFactor value in the.imd file). The units were converted from [\(W/(m^2*sr)\)] into [\(\mu W/(cm^2*nm*sr)\)] using nominal band pass widths of bands blue (68 nm), green (99 nm), red (71 nm), and NIR (114 nm) [37]. The IARR pre-processing was applied on the absolute radiance data to normalize the image and convert data to relative reflectance. The geometric correction was also implemented using topographic maps and drainage network of the studied area. The pre-processing of Landsat-8 includes dark subtract correction to reduce the atmospheric effects. The downloaded Landsat-8 data was pre-georeferenced to UTM zone 40-North projection. While pre-processing was applied to the data, the alteration minerals were enhanced by implementation of image-based algorithm including band ratioing and principal component analysis (PCA). Since Landsat-8 and
Quickbird are multi-spectral sensors containing few bands, it is better to use image-based algorithms instead of spectral-based ones such as linear spectral unmixing (LSU), matched filtering (MF), and mixture-tuned matched filtering (MTMF). In order to acquire information about alteration minerals and mineralogical characteristics at the Babbiduiyeh area, a field reconnaissance was carried out, and 50 samples were collected through a systematic sampling from fresh and representative hydrothermally altered rocks using a Global Positioning System (GPS) (Figure 1C). The mineralogy of samples was studied using XRD, spectroscopic studies, and transmitted reflected light microscopy. The ENVI 5.1, ArcGIS 10.3, PimaView, ViewSpecpro, and SAMS softwares were used for pre-processing and processing the satellite images and recognizing the field spectra.

4. Results and discussions
4.1. Mineralogical studies
The mineralogy of samples was studied using XRD, transmitted/reflected light microscopy, and spectroscopic studies. According to the major, minor, and trace phases of twenty one XRD analyses from field samples, four alteration zones including phyllic (13 samples), albitization (3 samples), silicic-phyllic (3 samples), and gossan (2 samples) were identified. Spectra of the samples were measured using an analytical spectral device (ASD) field-portable FieldSpec3® spectro-radiometer, which measures reflectance spectrum across a spectral range of 325–2500 nm with 10 nm individual band width. Multiple spectra were acquired from both fresh and weathered facets of the samples taking into account that they are mixtures of different minerals, and less abundant minerals may not be observed in one single measurement [7]. Approximately 138 spectra were measured from 50 samples, and were analyzed to characterize the hydrothermal and supergene alteration minerals using the Pima view V3.1 and ENVI 5.1 softwares. Furthermore, an automated mineral identification was implemented by creating a reference spectral library through PIMA View v3.1 to simplify and speed-up the process of spectral analysis. Since significant minerals such as quartz, feldspar, pyrite, and chalcopyrite do not have characteristic absorption features in the VNIR-SWIR region, extraction of mineral information was focused on the main hydrothermal alteration products, which could be recognized by spectroscopic analysis. The identified minerals include muscovite, illite, montmorillonite, goethite, hematite, Jarosite, chlorite, and epidote (Figure 3). According to the particular characteristics of spectra in VNIR-SWIR regions, phyllic-oxide, phyllic, propylitic, argillie, and oxide alteration (gossan) zones could be distinguished.

Figure 3. Different types of alteration minerals identified by spectroscopic studies A) argillie B) phyllic C) propylitic D) gossan.
Microscopic examination revealed that most of the samples essentially consisted of secondary minerals including sericite, kaolinite, epidote, chlorite, calcite, jarosite, goethite, and hematite formed at a later time through processes such as hydrothermal alteration and weathering. Furthermore, quartz, plagioclase, alkali feldspar, amphibolite, pyroxene, chalcopyrite, pyrite, titanum oxide, and magnetite are also observed in the samples (Figure 4). However, since the clay minerals were composed of very fine-grained particles in the majority of samples, it was very difficult to identify them correctly under microscopic studies.

### 4.2. Processing satellite images

#### 4.2.1. Band ratio

Band ratio is a multi-spectral image processing method that divides the digital number of one band to another one. Same surface minerals due to trend, topographic slope, shadow or seasonal changes in intensity and brightness angle of sun light indicate different radiance values [38]. These distorting effects cause errors in interpretation of supervisor and have misleading results. The band ratio process could minimize these environmental effects. Spectral band rationing enhances compositional information, while suppressing other types of information about earth’s surface. This method is very useful for highlighting certain features or materials that cannot be seen in the raw bands. Band ratio transformation is useful for a qualitative detection of hydrothermal alteration minerals [39, 40].

The argillic and phyllic zones are characterized by muscovite/illite and kaolinite. These minerals exhibit an intense Al-OH spectral absorption and reflectance features typically centered at 2.20 and 1.65 μm, respectively. According to the resample spectra of these minerals into Quickbird and Landsat-8, these absorption and reflectance features coincided with the bands 7 and 6 of Landsat-8, respectively (Figure 5). However, due to the lack of adequate bands in this region at Quickbird, it is not possible to identify clay minerals through this sensor. As a result, the band

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**Figure 4.** Microscopic thin sections of some samples observed in crossed polarized light A) Plagioclase phenocrysts are pervasively replaced by muscovites B) Veinlet of calcite with iron oxide C) Hornblane phenocrysts are replaced by chlorite and epidote D) Clinopyroxen phenocrysts are replaced by chlorite and epidote.
ratio of 6/7 of Landsat-8 was used to enhance the kaolinite, muscovite, and clay minerals. The discriminated areas were enhanced as bright pixels and were corresponded to the results obtained from the XRD and spectroscopic studies (Figures 5B and 5C). The VNIR bands contain chief information about the absorption features related to transition metals (e.g. Fe$^{2+}$, Fe$^{3+}$) within the oxide zone including jarosite, goethite, and hematite. These minerals exhibit prominent spectral absorption and reflectance features in the blue and red regions, respectively (near 0.44 and 0.67 μm), which coincide with band 2 of Landsat-8 and band 1 of Quickbird (Table 3, Figure 5). Therefore, it is possible to identify oxide minerals through Landsat-8 and Quickbird by applying the image-based techniques such as band ratio. The ratio 4/2 (Figures 6E and 6F) of Landsat-8 is useful for enhancing iron oxides because it has absorption in the blue and high reflectance in the red region. VNIR bands of Landsat-8 and Quickbird are nearly equivalent (Table 3), so that bands 4 and 2 of Landsat-8 are equal to bands 3 and 1 of Quickbird, respectively. Therefore, the band ratios 4/2 and 3/1 for Landsat-8 and Quickbird, respectively, are used for identification of the oxide zone or gossan. Comparison of the Quickbird and Landsat-8 results obtained from these band ratios revealed that the discriminated areas coincided (Figures 6E, 6F, 6H, and 6I). A higher spectral resolution in Landsat-8 facilitates the discrimination of more features such as clay and oxide minerals. Although it is not possible to discriminate clay minerals in Quickbird due to the lack of SWIR bands, a higher spatial resolution of Quickbird permits an accurate location of the features.

4.3. Principle component analysis (PCA)

PCA is a powerful statistical technique used to suppress irradiance effects and reduce the data redundancy by transforming the original data onto new principal component axes. Therefore, this technique produces an uncorrelated image, which has a much higher contrast than the original bands. PCA can be applied to the multivariate datasets such as multi-spectral remote sensing images, with the purpose of extracting specific spectral responses, as in the case of hydrothermal alteration minerals. This technique relies on establishing the relationship between the spectral responses of target materials such as alteration minerals. In this algorithm, numeric values could be extracted from the eigenvector matrix and used to calculate the principal component (PC) images. Using this relationship, it is possible to determine which PCs contain the spectral information of the target minerals by investigation of digital numbers (DNs) of pixels or eigenvector matrix containing high (bright) or low (dark) values [41]. The result of this method was uncorrelated PC bands equal to the number of bands. In order to determine which image contains information related to the spectral signatures of specific alteration and oxide minerals, eigenvector loadings in each PC image were examined. The PC image with moderate-to-high eigenvector loadings for diagnostic absorptive and reflective bands of the argillic-phyllic and oxide/hydroxide minerals were considered as the representative image for these minerals.

According to the spectral characteristics of iron oxide/hydroxide minerals (Figure 5A), bands 3 and 1 of Quickbird and bands 4 and 2 of Landsat-8 that correspond to the reflective and absorptive bands, can discriminate areas rich in these minerals. In addition, the reflective and absorptive bands 6 and 7 of Landsat-8 are suitable for detecting alteration minerals such as sericite, kaolinite, chlorite, and epidote. If the loading of the absorptive and reflective bands are negative and positive, respectively, in sign, the target area is shown by bright pixels; conversely, if the loading of the absorptive and reflective bands are positive and negative, the area is shown by dark pixels.

Tables 4 and 5 list the eigenvector matrix for the PCA results on the Quickbird and Landsat-8 bands. According to the eigenvector loading of Landsat-8, PC4 is appropriate for discrimination of alteration minerals as bright pixels due to the high positive loading in band 6 and high negative loading in band 7 (Figures 7B and 7C, and Table 4). PC5 for Landsat-8 could highlight areas of oxide/hydroxide minerals because it contains higher negative loadings for bands 1 and 2 and positive loading for band 4 (Table 4). The PC images are shown in Figures 7E and 7F. Investigation of the eigenvector loadings for Quickbird show that PC3 contains high loading for bands 1 and 3 in positive and negative signs, which is ideal for discrimination of oxide/hydroxide minerals or gossan areas as dark pixels (Table 5). In order to show the areas containing these minerals as bright pixels, the inverse of PC3 was obtained (Figures 7H and 7I).
Table 3. Comparison of VNIR bands of Quickbird and Landsat-8.

<table>
<thead>
<tr>
<th>VNIR bands of Landsat-8</th>
<th>VNIR bands of Quickbird image</th>
</tr>
</thead>
<tbody>
<tr>
<td>B2: 450-515 nm</td>
<td>B1: 450-520 nm</td>
</tr>
<tr>
<td>B3: 525-600 nm</td>
<td>B2: 520-600 nm</td>
</tr>
<tr>
<td>B4: 630-680 nm</td>
<td>B3: 630-690 nm</td>
</tr>
<tr>
<td>B5: 845-885 nm</td>
<td>B4: 760-900 nm</td>
</tr>
</tbody>
</table>

Figure 5. Spectral plots of some minerals extracted from USGS spectral library and resample to band centers of Landsat-8 and Quickbird. A) Some iron oxide/hydroxide/sulfate minerals B) some clay minerals.

Figure 6. A) ASD spectral plot of a sample containing illite (fresh facet) and goethite (weathered facet) and its location on image B) Band ratio 6/7 of Landsat-8 compared with spectroscopic studies and XRD, respectively. ASD spectral plot of a sample contains goethite and its location on image C) Band ratio 4/2 of Landsat-8 compared with spectroscopic studies and XRD, respectively. G) ASD spectral plot of a sample containing illite and its location on image H) Band ratio 3/1 of Quickbird compared with spectroscopic studies and XRD, respectively.
Table 4. Eigenvector matrix for PCA results of Landsat-8.

<table>
<thead>
<tr>
<th>Band</th>
<th>Band 1</th>
<th>Band 2</th>
<th>Band 3</th>
<th>Band 4</th>
<th>Band 5</th>
<th>Band 6</th>
<th>Band 7</th>
</tr>
</thead>
<tbody>
<tr>
<td>PC1</td>
<td>0.105</td>
<td>0.138</td>
<td>0.231</td>
<td>0.337</td>
<td>0.459</td>
<td>0.609</td>
<td>0.470</td>
</tr>
<tr>
<td>PC2</td>
<td>-0.745</td>
<td>-0.103</td>
<td>-0.117</td>
<td>-0.214</td>
<td>0.865</td>
<td>-0.148</td>
<td>-0.393</td>
</tr>
<tr>
<td>PC3</td>
<td>-0.345</td>
<td>-0.393</td>
<td>-0.410</td>
<td>-0.533</td>
<td>-0.068</td>
<td>0.448</td>
<td>0.262</td>
</tr>
<tr>
<td>PC4</td>
<td>-0.084</td>
<td>-0.075</td>
<td>0.110</td>
<td>0.189</td>
<td>-0.189</td>
<td>0.598</td>
<td>-0.739</td>
</tr>
<tr>
<td>PC5</td>
<td>-0.403</td>
<td>-0.491</td>
<td>-0.264</td>
<td>0.695</td>
<td>0.0424</td>
<td>-0.190</td>
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</tr>
<tr>
<td>PC6</td>
<td>0.490</td>
<td>0.244</td>
<td>-0.806</td>
<td>0.186</td>
<td>-0.001</td>
<td>0.111</td>
<td>-0.062</td>
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<tr>
<td>PC7</td>
<td>-0.675</td>
<td>0.713</td>
<td>-0.182</td>
<td>0.045</td>
<td>0.008</td>
<td>0.002</td>
<td>-0.012</td>
</tr>
</tbody>
</table>

Table 5. Eigenvector matrix for PCA results of Quickbird.

<table>
<thead>
<tr>
<th>Band</th>
<th>Band 1</th>
<th>Band 2</th>
<th>Band 3</th>
<th>Band 4</th>
</tr>
</thead>
<tbody>
<tr>
<td>PC1</td>
<td>0.366</td>
<td>0.475</td>
<td>0.581</td>
<td>0.550</td>
</tr>
<tr>
<td>PC2</td>
<td>0.428</td>
<td>0.345</td>
<td>0.213</td>
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</tr>
<tr>
<td>PC3</td>
<td>0.656</td>
<td>0.123</td>
<td>-0.714</td>
<td>0.211</td>
</tr>
<tr>
<td>PC4</td>
<td>0.503</td>
<td>0.800</td>
<td>-0.327</td>
<td>-0.011</td>
</tr>
</tbody>
</table>

Figure 7. A) ASD spectral plot for a sample containing illite (fresh facet) and goethite (weathered facet) and its location on image B C) PC4 of Landsat-8 compared with spectroscopic studies and XRD, respectively D) ASD spectral plot for a sample containing goethite and its location on image E F) PC5 of Landsat-8 compared with spectroscopic studies and XRD, respectively G) ASD spectral plot of a sample containing illite and its location on image H I) Inverse PC3 of Quickbird compared with spectroscopic studies and XRD, respectively.
5. Accuracy assessment

The discriminated altered areas were verified by field surveys, geological maps of the studied area, and laboratory analyses such as XRD, spectroscopic studies, and transmitted and reflected light microscopy. There is a distinct link between image processing results and analysis data in the Babbiduyeh area. The results of spectroscopic studies, which revealed the existence of goethite, hematite, muscovite, and illite in the samples, correspond to the band ratio and PCA results as well (Figures 6 and 7). According to the spectroscopic studies by ASD fieldSpec®, illite, muscovite, and kaolinite were observed in the argillic-phyllic zones and hematite, jarosite, goethite in the oxide zones. In some samples, Jarosite, goethite, and hematite were discriminated in the weathered surface, while illite and muscovite were observed in fresh surface (Figures 6 and 7 (A, D, and G)). These samples are normally from the supergen zones, where oxide minerals lie on phyllic and argillic alteration. The XRD results show the existence of quartz, muscovite, illite, albite, and orthoclase in the major phase, and goethite and jarosite in the minor phase (Figures 6 and 7 (C, F, and I)).

According to the geological map of the studied area, the discriminated areas coincided with the riodacite, granodiorite, and granite rocks. Thin sections of these rocks showed existence of Fe-oxide, chalcopyrite, pyrite, and titanium oxide.

6. Conclusions

In this research work, the alteration zones were mapped by applying the image processing methods such as band ratioing and PCA on Quickbird and Landsat-8. The results obtained were checked through field investigations so that the locations of ground control and sampling points were designed based on the enhanced alteration from image processing. Principal component analysis and band ratioing on Landsat-8 showed a good capability in highlighting the distribution of alteration and oxide-hydroxide mineral assemblages present in the alteration zones. Although the Quickbird data does not contain the SWIR band and could not discriminate alteration minerals, sufficient bands in VNIR can be used in mapping the distribution of iron oxide/hydroxide minerals. There is a general agreement between the results obtained from ASD and XRD analyses and petrographic studies. The combined spectroscopic measurements, XRD analysis, and petrographic studies revealed mineral assemblages typical of phyllic, phyllic-supergen, propylitic, argillic, and oxide alteration (gossan). Analyses of field samples including observations of thin sections, XRD analysis, and spectral measurements revealed that phyllic was most intensive from other hydrothermal alterations in the Babbiduyeh area. Since this type of alteration is usually directly associated with areas of copper mineralization, discrimination of phyllic-argillic alteration associated with oxide minerals could be beneficial in discrimination of PCD. In general, image processing of Quickbird, Landsat-8, and field samples revealed the effect of gossans on hydrothermal alteration zones, which could be useful in reconnaissance studies for porphyry copper identification.

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نقشه برداری دگرسانی گرمایی با استفاده از داده لندست-8 و کوئیک برد. مطالعه موردی: باب بیدوئیه، استان کرمان، ایران

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چکیده

در این پژوهش تمرکز ما بر روی مجموعه داده لندست-8 و کوئیک برای نقشه برداری دگرسانی گرمایی و گوسان در شناسایی کانه زایی مس پورفیری باب بیدوئیه است. این ناحیه در کمپلکس آتشفشانی-رسوبی ایران مرکزی جایی که ذخایر بزرگی مانند سرچشمه و چندین رخنمون متعدد مس وجود دارد قرار گرفته است. با اجرای آنالیزهای مؤلفه اصلی و نسبت باندی بر روی مجموعه داده تصویر لندست-8 و کوئیک برد، زونهای دگرسانی گرمایی شده. نتایج پردازش تصویر با استفاده از مطالعات میدانی بهره برانش اشعه ایکس (XRD)، نقاط نازک سیگنال‌های طیفی و مطالعات طیف‌سنجی از نمونه‌های صخره‌ای و نیز نقاط میدانی، نشان داده‌های طیفی نمونه‌ها با بررسی بصری و نرم‌افزارهای، PIMAView، SAMS و ViewSpecpro تجزیه و تحلیل شدند. نتایج این پژوهش نشان داد که دگرسانی‌های طیفی، آلایندر، XRD و مطالعات سنگ‌شناسی نشان داده که کانه‌ها نمونه‌های بارزی از دگرسانی‌های فیلیک-فیلیک-فلزی-فلزی-فلزی-فلزی (Alpine) و گوسان هستند. نتایج به دست آمده از پردازش تصویر و آلایندر نمونه‌های صخراهای تأثیر اکسیدهای و هیدروکسیدهای آهن را و سطح دگرسانی‌های فیلیک و آزرلیک نشان داد. برابر با سطح دگرسانی فیلیک و آزرلیک نشان داد که ناحیه کانه که گوسان بارز شده در کوئیک برد با نواحی دگرسانی فیلیک و آزرلیک مطابقت دارد.

کلمات کلیدی: نسبت باندی، دگرسانی گرمایی، لندست-8، آلایندر، مطالعات نمونه‌های اصلی، کوئیک برد.